







A new model of shoaling and breaking waves

1D solitary wave on a mild beach

Maria Kazakova, G. L. Richard, A. Duran



Modelling of surf zone propagation

Main issues

- Dispersive effects (dependence of wave phase velocity on its frequency)
- Energy dissipation
- Turbulent structures generation

Computational resources

- ► Navier-Stokes + BC : too expensive
 - Asymptotic models are used
- Depth-averaged equation



Overview

- ▶ Past & recent advances on wave breaking modelling : $NSWE \longrightarrow Boussinesq\,type \longrightarrow Shear\,flows$
- Model derivation key points
- Numerical realization
- Convergence of algorithm
- Experimental data comparison
 - Set up
 - Trial tests
 - Empirical laws
- Conclusions & Perspectives



State of the art

Advances on wave breaking modelling

Depth-average Euler equation $\longrightarrow \mu = \frac{h}{L}$ (SW), $\varepsilon = \frac{a}{h}$ (NL) NSWE BTGNSWPressure Hydrostatic Not hydrostatic Not hydrostatic NL small no assump. no assump. Shocks Х \times

Dissipation(?)

- * Artificial terms [Zelt, Kennedy, Cienfuegos,...]
- * Switching between model [Bonneton, Marche, Kazolea...]



State of the art

Recent advances, goals for the research

- G. Nwogu, **1996**, Y. Zhang *et al*, **2014** * Empirically add *TKE* equation.
- R. Briganti *et al.*, **2004**, S. L. Gavrilyuk *et al.*, **2016** * Use effects of vertical variations of the flow.

Research aims:

* To derive unified model capable to describe wave breaking
 ** Take into account turbulence structure explicitly
 ** Use a shear flow hypothesis



Filtered decomposition

Model derivation

Turbulence : Filtered decomposition

Filtering velocity field decomposition:

$$oldsymbol{v} = \overline{oldsymbol{v}} + oldsymbol{v}^r$$

(Navier-Stokes) gives

- * Filtered continuity equation
- * Filtered momentum equation
- * Decomposition of stress tensor
 - \rightarrow isotropic and anisotropic part



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Model derivation

Turbulence : Filtered decomposition

Continuity equation: div $\overline{\boldsymbol{v}} = 0$,

 $\text{Momentum equation: } \frac{\partial \overline{\boldsymbol{v}}}{\partial t} + \operatorname{\mathbf{div}}\left(\overline{\boldsymbol{v}\otimes\boldsymbol{v}}\right) = \boldsymbol{g} - \frac{1}{\rho}\operatorname{\mathbf{grad}}\overline{p} + \nu\Delta\overline{\boldsymbol{v}}$

 $\overline{v \otimes v}$ should be defined!

$$oldsymbol{\sigma}^r = -
ho\left(\overline{oldsymbol{v}\otimesoldsymbol{v}}-\overline{oldsymbol{v}}\otimes\overline{oldsymbol{v}}
ight) = -rac{2}{3}
ho k^r \mathbf{I} + \mathbf{A}^r$$

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Turbulent viscosity hypothesis

$$\mathbf{A}^{r} = 2\rho\nu_{T}\overline{D} = \rho\nu_{T} \left[\mathbf{grad} \ \overline{\boldsymbol{v}} + (\mathbf{grad} \ \overline{\boldsymbol{v}})^{T} + \text{Equation for kinetic energy}\right]$$

Model derivation

2D case

The filtered mass conservation

$$\frac{\partial \overline{u}}{\partial x} + \frac{\partial \overline{w}}{\partial z} = 0.$$

Momentum Ox balance equation :

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$$\frac{\partial \overline{u}}{\partial t} + \frac{\partial \overline{u}^2}{\partial x} + \frac{\partial \overline{u}\overline{w}}{\partial z} = -\frac{1}{\rho}\frac{\partial p}{\partial x} + \frac{1}{\rho}\left(\frac{\partial A_{xx}^r}{\partial x} + \frac{\partial A_{xz}^r}{\partial z}\right) + \nu\left(\frac{\partial^2 u}{\partial x^2} + \frac{\partial^2 u}{\partial z^2}\right)$$

Oz:

$$\frac{\partial \overline{w}}{\partial t} + \frac{\partial \overline{u}\overline{w}}{\partial x} + \frac{\partial \overline{w}^2}{\partial z} = -g - \frac{1}{\rho} \frac{\partial p}{\partial z} + \frac{1}{\rho} \left(\frac{\partial A_{xz}^r}{\partial x} + \frac{\partial A_{zz}^r}{\partial z} \right) + \nu \left(\frac{\partial^2 w}{\partial x^2} + \frac{\partial^2 w}{\partial z^2} \right)$$

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Model derivation

2D case

Equation for the kinetic energy :

$$\begin{split} &\frac{\partial}{\partial t} \left(\frac{\overline{u}^2}{2} + \frac{\overline{w}^2}{2} \right) \\ &+ \frac{\partial}{\partial x} \left[\overline{u} \left(\frac{\overline{u}^2}{2} + \frac{\overline{w}^2}{2} + gz \right) + \frac{pu}{\rho} - \frac{A_{xx}^r u}{\rho} - \frac{A_{xz}^r w}{\rho} - \frac{\tau_{xx} u}{\rho} - \frac{\tau_{xz} w}{\rho} \right] \\ &+ \frac{\partial}{\partial z} \left[\overline{w} \left(\frac{\overline{u}^2}{2} + \frac{\overline{w}^2}{2} + gz \right) + \frac{pw}{\rho} - \frac{A_{xz}^r u}{\rho} - \frac{A_{zz}^r w}{\rho} - \frac{\tau_{xz} u}{\rho} - \frac{\tau_{zz} w}{\rho} \right] \\ &= -\varepsilon_f - P^r. \end{split}$$

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Dissipation :

* ε_f – viscous dissipation due to filtered velocity field * P^r – energy transfer from large scale to small scale

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Virtual enstrophy : breaking criteria

$$\overline{u}(t,x,z) = U(t,x) + u'(t,x,z), \ \varphi = \frac{<< u'^2 >>}{h^2}$$

Teshukov, **2007**, Richard, Gavrilyuk, **2012**,**2013**,**2015** Integration over the depth and hypothesis of mild topography give:

$$\begin{aligned} & \left(\begin{array}{c} \frac{\partial h}{\partial t} + \frac{\partial hU}{\partial x} = 0, \\ & \left(\frac{\partial hU}{\partial t} + \frac{\partial}{\partial x} \left(hU^2 + \frac{gh^2}{2} + h^3\varphi + \frac{h^2\ddot{h}}{3} \right) = \frac{\partial}{\partial x} \left(\frac{4}{R} h^3 \sqrt{\varphi} \frac{\partial U}{\partial x} \right) - ghb', \\ & \left(\begin{array}{c} \frac{\partial h\varphi}{\partial t} + \frac{\partial hU\varphi}{\partial x} = \frac{8h\sqrt{\varphi}}{R} \left(\frac{\partial U}{\partial x} \right)^2 - C_r h^3 \varphi^{3/2} \end{aligned} \right) \end{aligned}$$

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Hyperbolic & Elliptic stage separation

$$\begin{split} \frac{\partial h}{\partial t} &+ \frac{\partial (hU)}{\partial x} = 0, \\ \frac{\partial (hU)}{\partial t} &+ \frac{\partial}{\partial x} \left(hU^2 + \frac{gh^2}{2} + \boxed{\frac{h^2\ddot{h}}{3}} + h^3\varphi \right) = \frac{\partial}{\partial x} \left(\frac{4h^3\sqrt{\varphi}}{R} \frac{\partial U}{\partial x} \right) - b'hg, \\ \frac{\partial h\varphi}{\partial t} &+ \frac{\partial (h\varphi U)}{\partial x} = \frac{8h\sqrt{\varphi}}{R} \left(\frac{\partial U}{\partial x} \right)^2 - C_r h^3 \varphi^{3/2}, \end{split}$$



Hyperbolic & Elliptic stage separation

$$\begin{cases} \frac{\partial h}{\partial t} + \frac{\partial (hU)}{\partial x} = 0, \\ \frac{\partial (hK)}{\partial t} + \frac{\partial}{\partial x} \left(\frac{hKU}{R} + \frac{gh^2}{2} + \alpha + h^3 \varphi \right) = \frac{\partial}{\partial x} \left(\frac{4h^3 \sqrt{\varphi}}{R} \frac{\partial U}{\partial x} \right) - b'hg, \\ \frac{\partial h\varphi}{\partial t} + \frac{\partial (h\varphi U)}{\partial x} = \frac{8h\sqrt{\varphi}}{R} \left(\frac{\partial U}{\partial x} \right)^2 - C_r h^3 \varphi^{3/2}, \end{cases}$$

Change of variables $K = u + \frac{1}{3h} \nabla \left(h^2 \dot{h} \right)$

O. Le Métayer, S. Gavrilyuk, S. Hank, (2010) A numerical scheme for the Green–Naghdi model, JCP



Hyperbolic Stage $h,\,hK,\,h\varphi$

Hyperbolic stage \rightarrow Godunov's Type Scheme of 2^{nd} order

- $* 2^{nd}$ order : MUSCL Reconstruction
- * Riemann Solver : HLL
- * Time discretization : RK-2 (Heun's method)



Elliptic operator inversion

Elliptic stage \rightarrow System of linear equations

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Elliptic equation for velocity $hK = hU - \frac{1}{3}\nabla\left(h^2\frac{\partial hU}{\partial}\right) + \frac{1}{6}\nabla\left(hU\nabla(h^2)\right)$ In discrete form with three-diagonal matrix A

$$\begin{pmatrix} \vdots \\ hU \\ \vdots \end{pmatrix}_{i}^{n+1} = \begin{bmatrix} A\left(h,h^{2},\frac{\partial h^{2}}{\partial x},\frac{\partial^{2}h}{\partial x^{2}}\right)_{i}^{n+1} \end{bmatrix} \begin{pmatrix} \vdots \\ hK \\ \vdots \end{pmatrix}_{i}^{n+1}$$

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Well-balanced algorithm



E. Audusse et al, (2004) A fast and stable well-balanced scheme with hydrostatic reconstruction for shallow water flows, J Sci Comput, 25(6), 2050-2065.



Virtual enstrophy : breaking criteria $\forall \varepsilon < 0.05$, or at the initial wave propagation ($t < t^*$)

$$\begin{cases} \frac{\partial h}{\partial t} + \frac{\partial hU}{\partial x} = 0, \\ \frac{\partial hK}{\partial t} + \frac{\partial}{\partial x} \left(hUK + \frac{gh^2}{2} + \alpha \right) = -ghb', & \forall t < t^*, \forall x \in [0, L] : \\ \max_x \psi(t, x) < \psi_0, \\ \frac{\partial h\psi}{\partial t} + \frac{\partial(hU\psi)}{\partial x} = \frac{8h\sqrt{\psi}}{R} \left(\frac{\partial U}{\partial x}\right)^2 - \mathcal{D}, \end{cases}$$

Virtual enstrophy : breaking criteria

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$$\begin{cases} \frac{\partial h}{\partial t} + \frac{\partial hU}{\partial x} = 0, \\ \forall t > t_*, \quad \frac{\partial hK}{\partial t} + \frac{\partial}{\partial x} \left(hUK + \frac{gh^2}{2} + h^3\varphi + \alpha \right) = \frac{\partial}{\partial x} \left(\frac{4}{R} h^3 \sqrt{\varphi} \frac{\partial U}{\partial x} \right) \\ \forall x \in [0, L] \quad \frac{\partial h\psi}{\partial t} + \frac{\partial (hU\psi)}{\partial x} = \frac{8h\sqrt{\psi}}{R} \left(\frac{\partial U}{\partial x} \right)^2 - \mathcal{D}, \\ \forall t > t_*, \\ \forall x \in \bigcup_{\ell=t^*}^t \{ x : \max_x \psi(\ell, x) > \psi_0 \}, \quad \frac{\partial h\varphi}{\partial t} + \frac{\partial (hU\varphi)}{\partial x} = \frac{8h\sqrt{\varphi}}{R} \left(\frac{\partial U}{\partial x} \right)^2 - \mathcal{D}, \end{cases}$$

Soliton test & Convergence

$$h(x,t) = h_0 + \xi(x,t), \quad u(x,t) = c_0 \left(1 - h_0/h(x,t)\right)$$

$$\xi(x,t) = \frac{2a \left(Fr^2 - 1 - 3\tilde{\varphi}\right)}{Fr^2 - 1 - (3 + a^2)\tilde{\varphi} + (Fr^2 - 1 - (3 - a^2)\tilde{\varphi})\cosh(\kappa(x - c_0t - x_0))}$$

$$\kappa = \sqrt{\frac{3(Fr^2 - 1 - 3\tilde{\varphi})}{Fr^2}}, c_0 = \sqrt{g \left(h_0 + a + \tilde{\varphi}(3h_0 + a)\right)}$$



Soliton test & Convergence



x



900

Experimental Data Comparison

S. C. Hsiao *et al.*, **2008**,
$$tg\beta = 0.017$$



Experimental Data Comparison

S. C. Hsiao *et al.*, **2008**,
$$tg\beta = 0.017$$



Experimental Data Comparison

S. C. Hsiao *et al.*, **2008**, $tg\beta = 0.017$



Experimental Data Comparison





Experimental Data Comparison



Empirical Laws for breaking depth



Empirical Laws for breaking depth ()





CLE. Synolakis, J. Skjelbreia, Evolution of maximum amplitude of solitary waves on plane beaches,

J. Waterw. Port Coast, Ocean Eng. 119 (3), 323-342, (1993).

Empirical Laws for breaking depth ()

$$\begin{aligned} \frac{\eta_{max}}{h_b} &\sim \left(\frac{h_{loc}}{h_b}\right)^n\\ n &= -1/4 - \text{gradual shoaling, } n = -1 - \text{rapid shoaling}\\ n &= 4 - \text{rapid decay, } n = 1 - \text{zone of gradual decay} \end{aligned}$$
$$\psi_0 &= \frac{g}{h_0^*} \widetilde{\psi}_0, \widetilde{\psi}_0 = \begin{cases} \left(0.1 + \frac{0.031}{\varepsilon}\right), & \varepsilon > 0.05, \\ 0, & \varepsilon < 0.05, \end{cases} R = \begin{cases} 1.7, & \varepsilon > 0.05, \\ 6, & \varepsilon < 0.05, \end{cases}$$



Wave transformation over different slopes



Wave transformation over different slopes

Dependence of Reynolds number R on topography $tg \beta$ * $tg \beta = 0$ from stability of solitary wave $\varepsilon = 0.78$.



Conclusions & Perspectives

- Modelling
 - Dispersive model for shoaling and breaking waves is derived
 - Natural procedure for the breaking wave description is proposed
- Numerics
 - The well-balanced 1D code is developed
 - Comparison with exact solution, experimental data is performed
 - Numerical simulations are confirmed with empirical laws for breaking depth

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Purpose for the further investigation:

1D Non-uniform topography with dry zones (in prep. Richard, Duran) Cnoidal waves propagation

2D simulations (in prep. Richard, Duran, Fabrèges)